

Modeling Wave Propagation Using Shallow Water Equation

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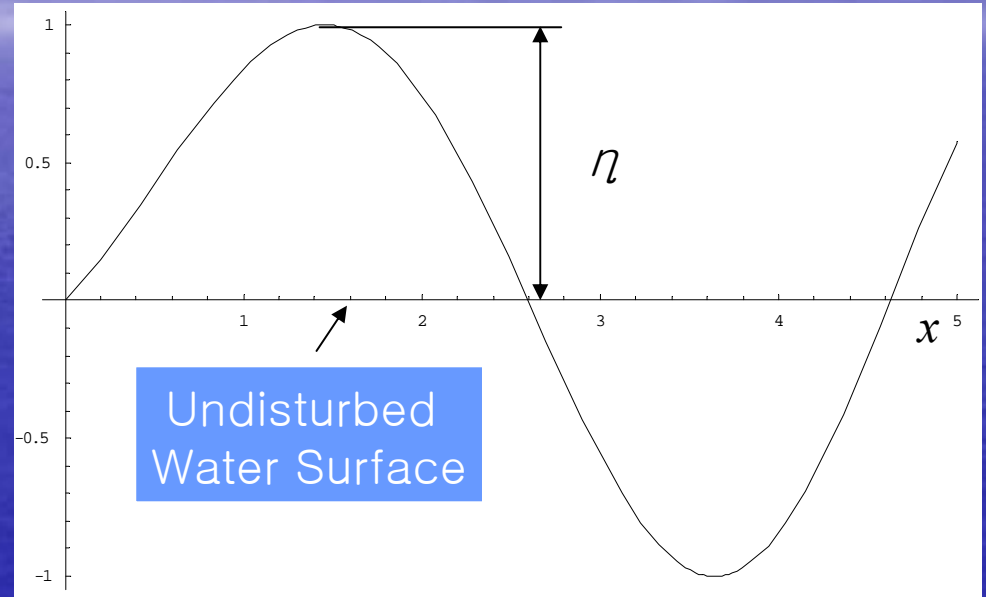
Characteristic of Shallow Water Eq.

- No pressure variance in vertical direction \rightarrow Long wavelength
- Mass/Momentum Conservation
- Inviscid fluid flow

Baby Shallow Water Eq. (1D)

$$\frac{\partial \eta}{\partial t} + \frac{\partial M}{\partial x} = 0$$

$$\frac{\partial M}{\partial t} + gD \frac{\partial \eta}{\partial x} = 0$$



η : water level relative to equilibrium level

D : Total thickness of water

M : depth averaged velocity in $+x$ direction* D

Numerical Scheme (I)

$$\frac{\partial \eta}{\partial t} + \frac{\partial M}{\partial x} = 0$$

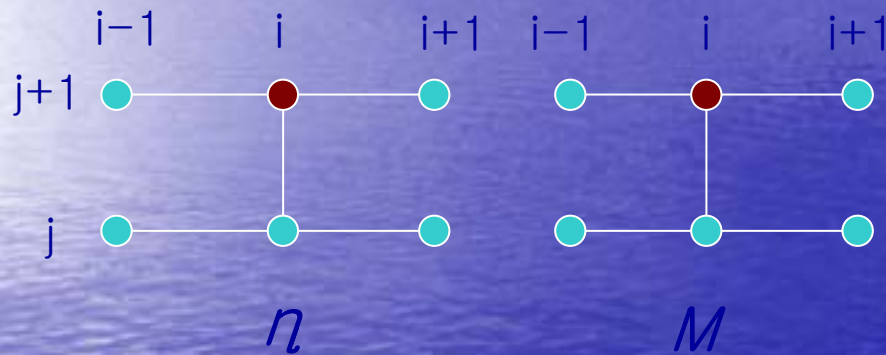
$$\frac{\partial M}{\partial t} + gD \frac{\partial \eta}{\partial x} = 0$$

- Finite Difference Method
 - 2nd order centered difference : $\partial/\partial x$
 - 2nd order centered difference : $\partial/\partial t$(Crank–Nicolson Method)
- Boundary condition
 - $M = 0$ for boundaries (no velocity)

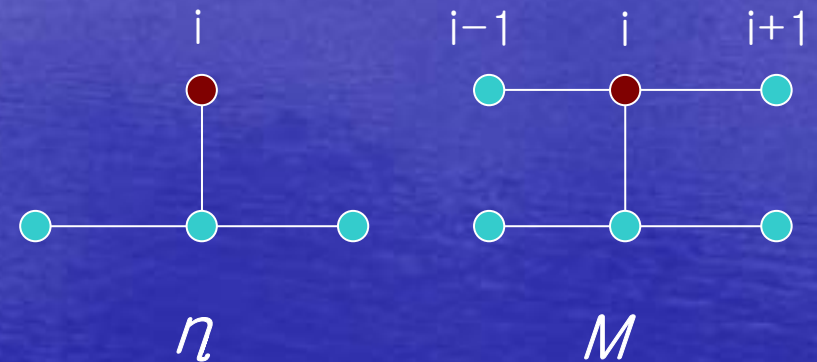
Numerical Scheme (II)

Implicit vs Explicit

Implicit



Semi-Explicit



$$\frac{\partial \eta}{\partial t} + \frac{\partial M}{\partial x} = 0$$

$$\frac{\partial M}{\partial t} + gD \frac{\partial \eta}{\partial x} = 0$$



Finite
Difference
Method

.....

Three Initial Conditions



Sine curve

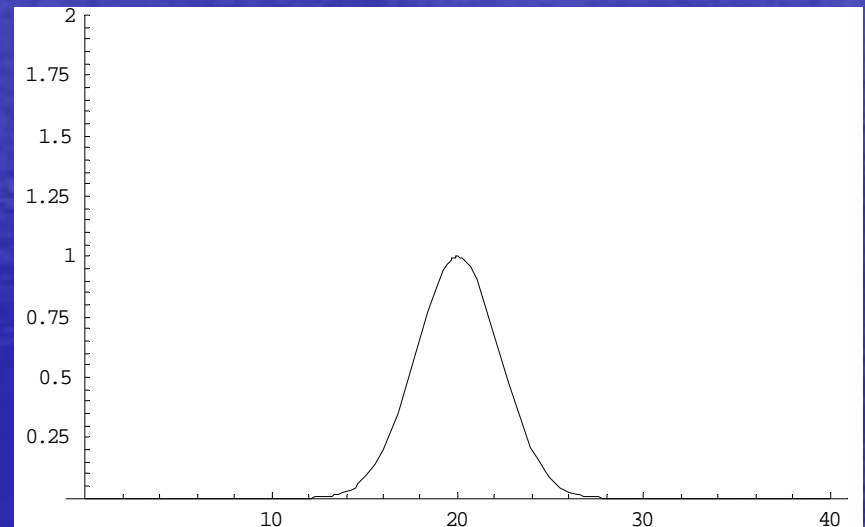
Application:

Tsunami wave approaching shore

- Has long wavelength (~ 100 km)
- small amplitude (~ 1 m)
- Travels at 430 mi/hr

- Initial Condition

– η
– M



A Gaussian

Uncoupled SWEq.

$$\frac{\partial^2 M}{\partial t^2} = g(\eta + h) \frac{\partial^2 M}{\partial x^2} - \frac{1}{\eta + h} \frac{\partial^2 M}{\partial t \partial x}$$

$$\frac{\partial^2 \eta}{\partial t^2} = g(\eta + h) \frac{\partial^2 \eta}{\partial x^2} + g \left(\frac{\partial \eta}{\partial x} \right)^2 + \frac{\partial h}{\partial x} \frac{\partial \eta}{\partial x}$$



Speed of Wave
Propagation

$$= \sqrt{g(\eta + h)} = \sqrt{gD}$$

$D = \eta + h$ (total thickness of water)

h : depth profile

Big Shallow Water Eq. (2D)

$$\frac{\partial \eta}{\partial t} + \frac{\partial M}{\partial x} + \frac{\partial N}{\partial y} = 0$$

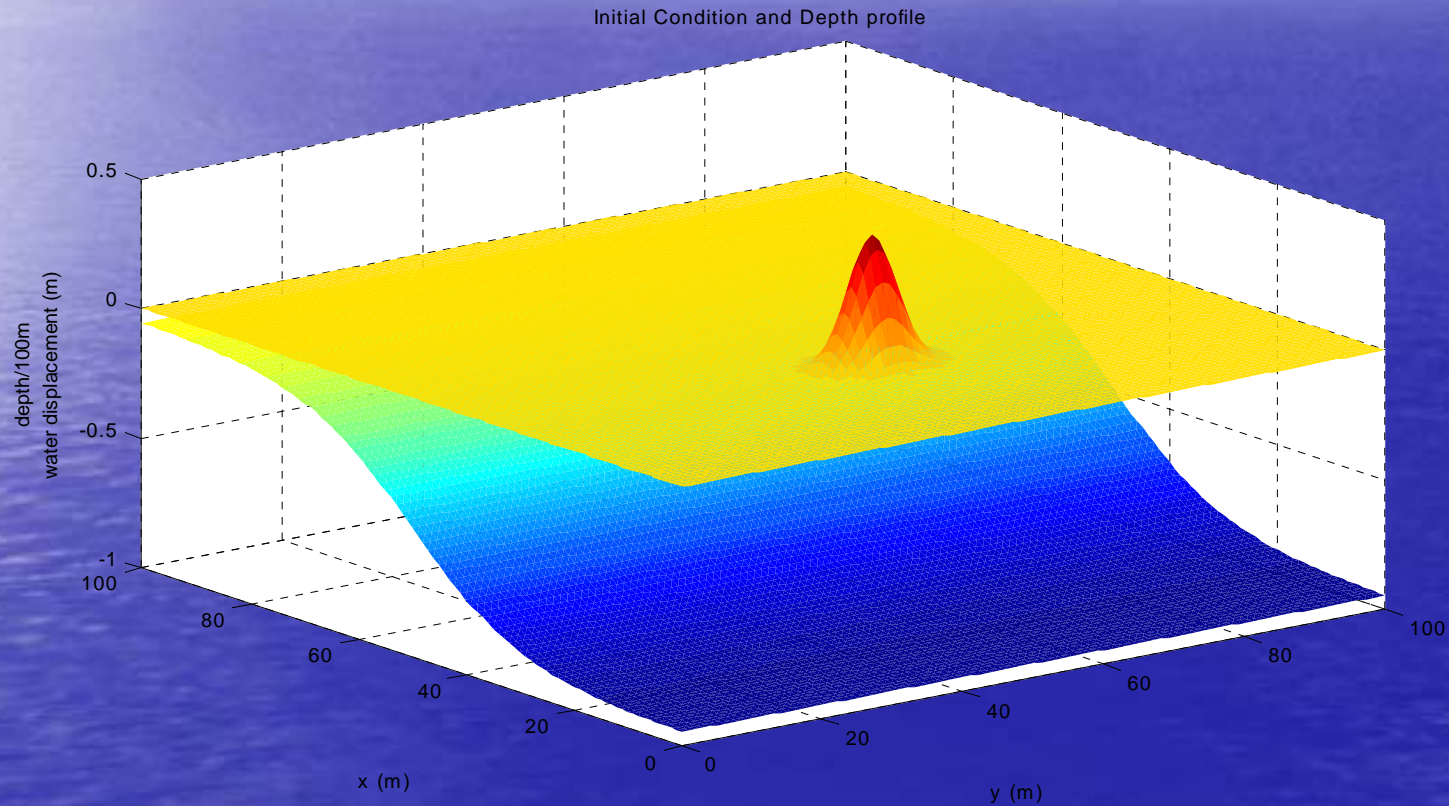
$$\frac{\partial M}{\partial t} + \frac{\partial}{\partial x} \left(\frac{M^2}{D} + \frac{1}{2} g \eta^2 \right) + \frac{\partial}{\partial y} \left(\frac{MN}{D} \right) + gh \frac{\partial \eta}{\partial x} = 0$$

$$\frac{\partial N}{\partial t} + \frac{\partial}{\partial x} \left(\frac{MN}{D} \right) + \frac{\partial}{\partial y} \left(\frac{N^2}{D} + \frac{1}{2} g \eta^2 \right) + gh \frac{\partial \eta}{\partial y} = 0$$

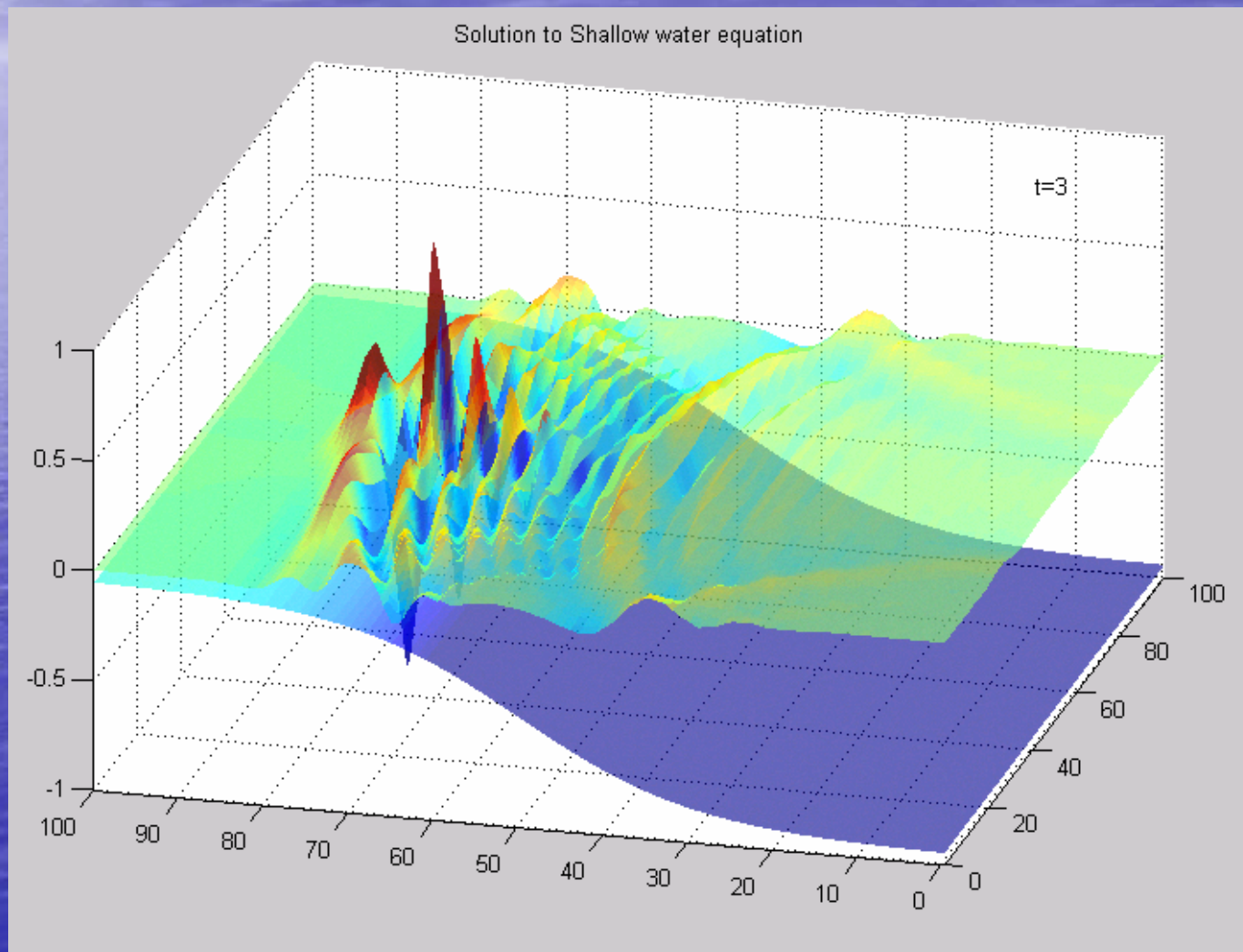


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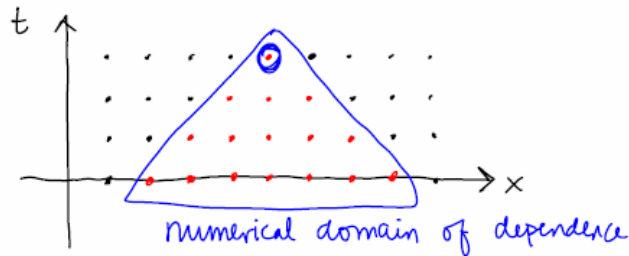
Simulation: Initial Condition / Depth profile



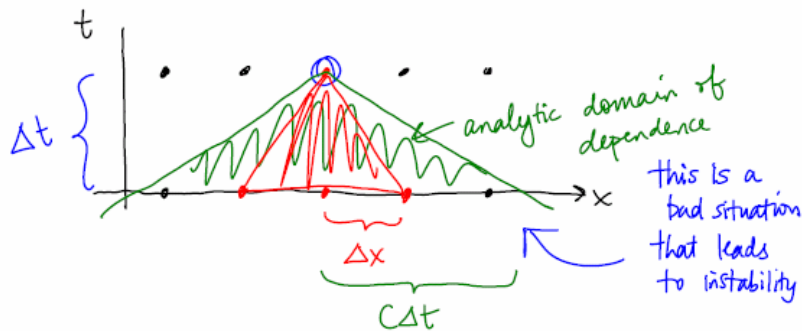
Numerical Stability



The Key



Imagine one time step taken:



$$u_{tt} = c^2 u_{xx} + \frac{c^2}{12} \Delta x^2 (1 - \alpha^2) u_{xxxx} + O(\Delta x^4, \Delta t^4)$$

so this is why the error seems less when $\alpha=1$

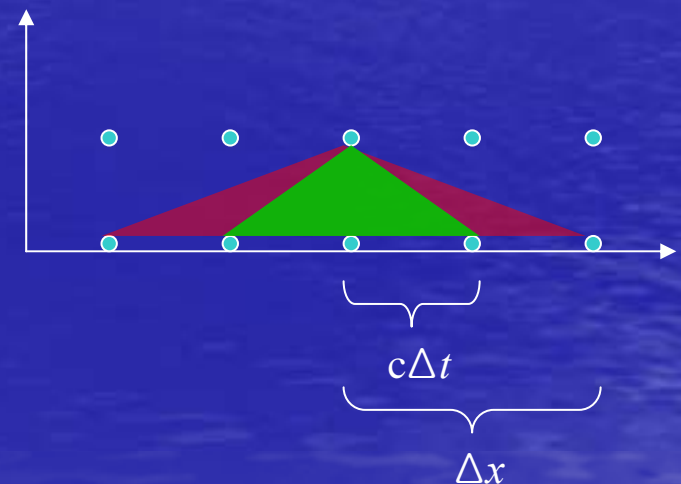
This term causes the numerical scheme to be dispersive (i.e., waves with different wavenumbers travel at different speeds)

- Analytic domain of dependence is determined by the speed of wave propagation

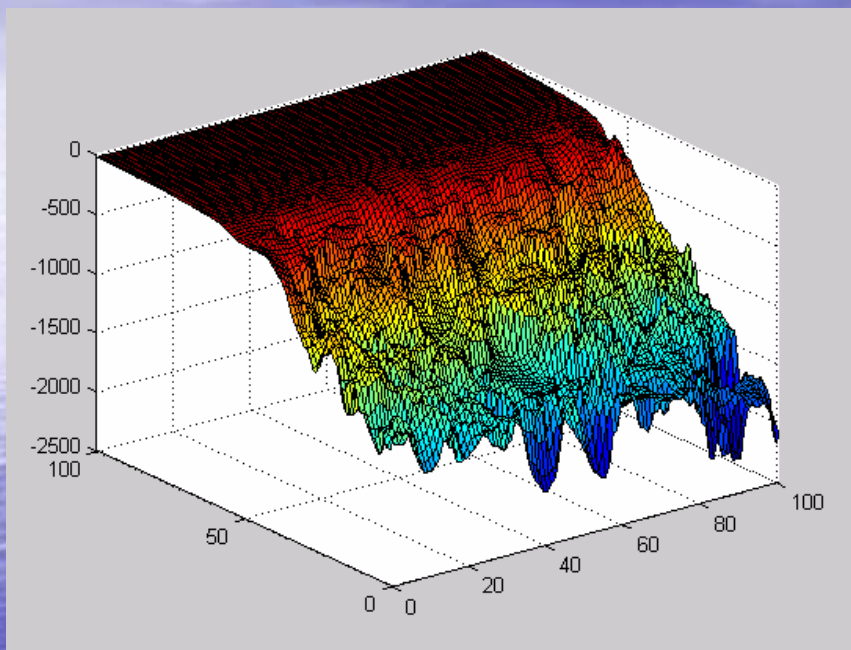
- Speed of Wave Propagation

$$c = \sqrt{g(\eta + h)} = \sqrt{gD}$$

- As $D \downarrow$, $c \downarrow$

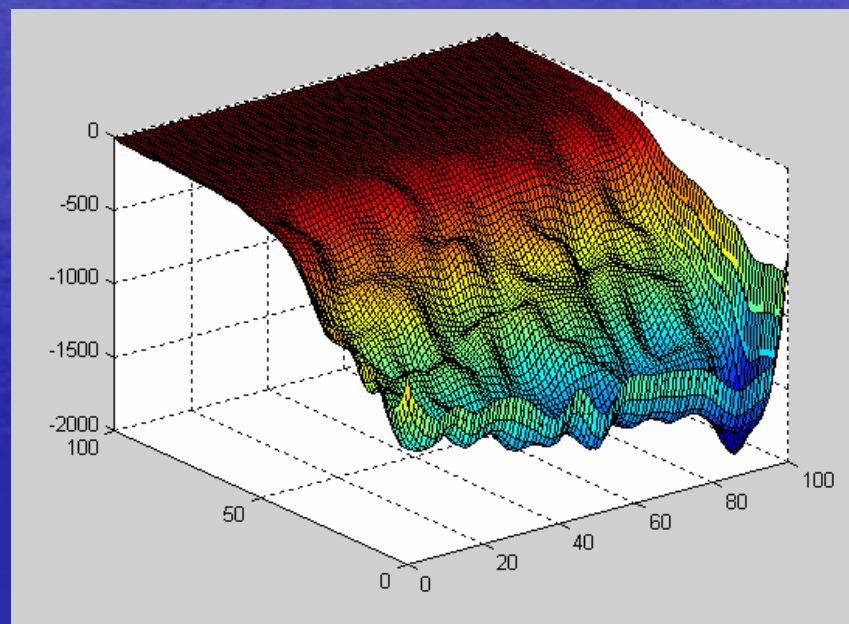


Gulf of Mexico



Take the convolution of each point with a Gaussian

convolution



Further Work

- Adaptive grid $\rightarrow \Delta x(h)$

This is would account for speed variation

- Effective Equation Analysis
- Flux Limiter

Acknowledgement

- Professor Yong
- Nick Alger